## THEORETICAL RESEARCH OF MOVEMENT MOUNTAIN STREAM OF FLOWS

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Mathematical model of motion of different type multi – phase media on inclined surface describing the current avalanche-like flows is constructed on the basis of averaging method. We consider turbulence models of different complexity with regard to theology of multi – phase media by means of which the parameters of turbulent transfer appearing flow equations may be given.

The structure of avalanche-like flows is may be represented in the form of granulated liquid that is a group of discrete solid particles dispersed in liquid in such a way that the particles of solid component are in contract with neighboring particles.

Landslips, mudflows, snow avalanches, flooding are dangerous natural processes and they create obstacles in constructing civil and industrial buildings, roads, developing lands and extracting mineral products. They are dangerous for human life. At present everywhere realization of measures in engineering defense of territories and objects is not always sensible and economically warranted. There fore, study of motion of landslips, mudflows, snow avalanches and prediction of development of these processes are very urgent and have practical value.

The requirement to mathematical models of motion of levanter – like flows is that the model must be constructed on the base of mechanics of composite media in the frames of macroscopic – phenomenological representations and simultaneously possess the properties of granulated medium (hard component), liquid and gas (water, air).

In describing motion of avalanche-like flow we use a model of multi- phase medium that "pours" down (motion arises from rest state). Variable steepness slope - that is long and wide in the form of variable length and depth channel, i.e. the effects related with interactions on lateral boundaries of the flow, are taken into account. The flow motion is subjected to the action of gravity force and friction force. This model may be constructed by various methods. In some cases for each mixture component we can write a system of equations and phase interactions take into account with the help of additional terms. In other cases we can approximately represent multi-phase flow as homogeneous one with averaged properties. The advantage of the last method is comparative simplicity of statement and solution of practical problems. Approximation degree of obtained results that in some cases are far from truth, is the defect of this method. In models considered up to now where turbulence of the flow is taken into account [1], the multi-phase property was ignored, if multi-phase property was taken into account, the turbulence was ignored [2]. It was assumed that for describing turbulence at a point it suffices to know only the scale of velocity, but stress components may be expressed by this scale by means of Kolmogorov-Prandtl relation for turbulent viscosity. However, in transferring components, the stresses equitably reflected by these relations, even if the transfer of velocity scale described. In complicated turbulent flows of multi-phase avalanche-like difference in evolution of different. Reynolds stress components should be consideration and transfer of the components must be taken properly count. To this end the models using transfer equations for

separate compo-turbulent stresses  ${U}_i \cdot {U}_j$  [2] must be worked out. These equations may be

exactly, but for obtaining a close system, modal relations must be introduced. The advantage of exact equations is that there automatically appear the king into account buoyancy forces, rotational motions and other factors and a they are very complicated, especially if they satisfy invariance and realize – requirements. Therefore, in spite of their great potential abilities they are not used in practice up to now. In motion of various character avalanche-like consideration of temperature change of medium is inevitable. Them express the dissipative turbulent motions are isotropic (local isotropy). Then express the dissipative term as follows:

$$\varepsilon_{ijj} = \frac{2}{3}\varepsilon\sigma_{ij}$$

 $\mathcal{E}$  is velocity of dissipation of kinetic energy  $K = \frac{1}{2} \overline{U_i \cdot U_j}$ , (scalar products name quantities

are denoted by dash). Equation for a general case is not cited here because of its awkwardness. Change of these quantities in horizontal direction may be determined by two- dimensional equations of mean flow for values averaged in depth. Grating three- dimensional equations in depth and assuming that pressure is hydrostatic, we can derive mean flow equations for quantities averaged. As a result we get: tinnitus equation

$$\frac{\partial h}{\partial t} + \frac{\partial}{\partial x} \left( h \overline{u} \right) + \frac{\partial}{\partial y} \left( h \overline{v} \right) = 0 \tag{1}$$

Equation of motion amount (in direction X)

$$\frac{\partial \overline{u}}{\partial t} + \overline{u} \frac{\partial \overline{u}}{\partial x} + \upsilon \frac{\partial \overline{\upsilon}}{\partial y} = -g \frac{\partial}{\partial x} (h + z_b) + \frac{1}{\rho h} \left( \frac{\partial (h \overline{\tau}_{xx})}{\partial x} + \frac{\partial (h \overline{\tau}_{xx})}{\partial y} \right) + \frac{\tau_{xx} - \tau_{bx}}{\rho h} \left( \frac{\partial \overline{F}}{\partial x} + \frac{\partial \overline{G}}{\partial y} \right)$$
(2)

Equation of motion amount (in direction Y)

$$\frac{\partial \overline{\upsilon}}{\partial t} + \overline{u} \frac{\partial \overline{u}}{\partial x} + \overline{\upsilon} \frac{\partial \overline{\upsilon}}{\partial y} = -g \frac{\partial}{\partial y} (h + z_b) + \frac{1}{\rho h} \left( \frac{\partial (h \overline{\tau}_{xy})}{\partial x} + \frac{\partial (h \overline{\tau}_{yy})}{\partial y} + \tau_{sy} - \tau_{by} + \frac{\partial \overline{F}}{\partial x} + \frac{\partial \overline{G}}{\partial y} \right)$$
(3)

Equation of medium temperature

$$\frac{\partial \overline{\Phi}}{\partial t} + \overline{u} \frac{\partial \overline{\Phi}}{\partial x} + \overline{\upsilon} \frac{\partial \overline{\Phi}}{\partial y} = \frac{1}{\rho h} \left( \frac{\partial (hJ_x)}{\partial x} + \frac{\partial (h\overline{J}_y)}{\partial y} + \frac{\partial \overline{P}}{\partial x} + \frac{\partial \overline{Q}}{\partial y} \right)$$
(4)

The last two terms in equations (2) - (4) are dispersing.

Here  $\overline{u}$  and  $\overline{v}$  h(t, x, y) is flow depth,  $\tau_s$ ,  $\tau_b$  are tangential stresses of surface and bottom, respectively,  $q_s$  is distributed heat flow through surface (heat flow from bottom may be assumed to be equal to zero in majority of cases), are denoted:

$$\overline{F} = \int_{z_b}^{z_b+h} \rho(u-\overline{u})(v-\overline{v})dz; \qquad \overline{G} = \int_{z_b}^{z_b+h} \rho(u-\overline{u})^2 dz;$$
$$\overline{P} = \int_{z_b}^{z_b+h} p(u-\overline{u})(v-\overline{v})dz; \qquad \overline{Q} = \int_{z_b}^{z_b+h} \rho(v-\overline{v})^2 dz;$$
$$\overline{\Phi} = \frac{1}{h} \int_{z_b}^{z_b+h} \Phi dz; \qquad U = \frac{1}{h} \int_{z_b}^{z_b+h} U dz; \qquad \overline{v} = \frac{1}{h} \int_{z_b}^{z_b+h} v dz;$$

here  $z_b$  are coordinates of inclined surface streamlined by multi- phase, multi- component flow. Now, turbulence models are necessary in order to determine depth averaged values of turbulent stresses  $\overline{\tau}_{ij}$  and depth averaged flows  $\overline{J}_i$  and also bottom tangential stresses (friction stresses)  $\tau_b$ . Flow shift at the surface (wind shift) and heat influx from the heat surface is determined by turbulent average flow of mixed mass. The quantities  $\tau_b$  and  $q_s$  are usually determined by means of simple empiric laws that contain friction coefficients and heat transfer. In the equations mentioned above, in many practice situations we can neglect terms containing turbulent stresses  $\overline{\tau}_{ij}$  averaged in depth in comparison with other terms, and effect of turbulence tells tells only on tangential bottom stresses  $\tau_b$ . In these cases only relation of bottom tangential stresses depth averaged velocities should be determined in model. To this end we use the relation derived from the friction law

$$\tau_{bx} = \rho \cdot c_f \overline{u} \sqrt{\overline{u}^2 + \overline{\upsilon}^2}; \qquad \tau_{by} = \rho \cdot c_f \overline{\upsilon} \sqrt{\overline{u}^2 + \overline{\upsilon}^2}; \qquad (5)$$

Where  $c_f$  is empirical friction coefficient dependent on the state of the given surface (the values for smooth and rough bottoms may be found in technical references). Unlike turbulent stresses, turbulent mass and heat flows  $\overline{J}_i$  always play an important role and require model assumptions. In equations (1) – (4) there are also so-called dispersive terms that appear because of vertical inhomogeneous of quantities related with mean flow. These terms have no connection with turbulence, they arise only because of depth averaging process and modeling of these terms is not turbulence modeling. All the turbulence models considered here, are based on turbulent viscosity (diffusion) concept. A great majority of investigations of such content assumptions such as steadiness of turbulent viscosity coefficients whose acceptance essentially simplifies the mathematical problem.

When turbulence is generated mainly in bottom and lateral areas (e.g. flow in riverbeds) depth averaged coefficient of turbulent diffusion is well described by the expression that contains dynamic velocity  $u_{\tau}$  and flow depth *h*:

$$\overline{R}_t = c \cdot \overline{u}_\tau \cdot h \tag{6}$$

Here c is an empiric constant depending on flow geometry ( $c \approx 0,135 \div 0,275$ ) and it changes depending on the ratio of flow width to its depth. Notice that in calculations of depth averaged quantities the depth average value of the turbulent diffusion coefficient  $\overline{R}_t$  belong to transfer on the whole not only to turbulent but also to dispersive one that arises because of heterogeneity of vertical distribution of velocities, temperatures and concentrations. Potentially universal model of avalanche like flows should allow consideration of back round outbursts to environment. Limit values of quantities typical for free are given on free boundaries. On rigid boundaries (lateral surface and button) immediately after viscous under layer it is given resultant velocity ( $u_{pez}$ ) allowing for boundary layer (of variable thickness) adhesion of heterogeneous media:

$$\left(u_{per}\right) = \frac{u_r}{\chi} \ln\left(y_0^+ \cdot E_0\right) \tag{7}$$

Where  $y_0 \cdot u_\tau / v$ ;  $\chi$  is the Carman constant (~0,4),  $E_0$  is roughness coefficient

 $(\approx 9 \div 17)$ , the point  $y_0$  is chosen so that  $y_0^+$  the dimensionless distance from the wall  $y_0^+$  is in the interval 30÷100;  $\nu$  is kinematics molecular viscosity coefficient.

In the paper we considered different degree complexity turbulence models with regard to reology of multi-phase madia by whose means we can give turbulent transfer parameters appearing in averaged flow equations. These models were worked out to investigate natural phenomena such as motion of avalanche- like flows and air dynamics of environment in nature catastrophes as floodings and etc. Elaboration of effective methods of solution of practical problems on the base of the suggested model will be represented in future papers of the author.

## Literature

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